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HYDROLOGIC MODEL OP GUAYANILLA BAY, PUERTO RICO

Michael A. Chartock

February 1980

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Center for Energy and Environment Research

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Mayaguez, P.R. 00708

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#### ACKNOWLEDGEMENT AND DISCLATHER

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## EXECUTIVE SUMMARY

Objectives and purpose: hydrologic model is developed to  
Account for the average water flow patterns in Guayanilla Bay  
as a guide to understand the function of the Bay, to identify  
controls over bay processes, to describe important data un-

certainties, and to inform decisions on the effective uses

of the bay.

Wind, tidal forces, industrial pumping and freshwater input are

The major external forces affecting the bay, although each of the bay's five major compartments are unique in the extent to which these forces dominate. Precipitation, evaporation and ship activity affect the water budget to a much lesser degree. Bottom topography or bathymetry of the bay is the most critical feature controlling the flux of water and the signature characteristics of the bay compartments.

Biological processes control the rate of exchange between several compartments. Mangroves are important in the thermal cove and Southeast Bay compartments. Seagrasses are important in the Western Bay, Southeastern Bay and Central bay as they affect

Sill depth and wind drift,

Wind drift and equilibrium (windsurface) return flows are both the largest flows and those associated with the greatest uncertainty. Based on this study a research effort to characterize wind drift flow rates across shallow bay sills, together with

an evaluation of equilibrating return flows in deeper channels would be the most beneficial studies for predicting the physical behavior of Guayanilla Bay.



Management options affecting the bay can influence the bay's

Productivity, use as a port, and the characteristics of water

used for power plant cooling. Wind, geomorphology, and intake

and discharge location have the greatest control over power plant

Cooling water intake temperature, sills between bay compartments

and the freshwater from surface and groundwater of the Yauco and

Guayanilla River watersheds have the greatest control over

biological productivity (from a hydrologic standpoint). Currents,

and biological communities could be managed to control sedimentation

Fates and stabilize bottom topography. Thus, management decisions

in development of the bay can be informed by a hydrologic model

to sustain and enhance productive uses in an efficient manner.

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## Resumes

Objetivos y propósito: se desarrolló un modelo hidrológico que con-  
sidera el patrón de flujo promedio de las aguas de la zona de  
Guayanilla, el mismo sirve de guía para comprender el funciona-  
miento de la bahía, identificar lo que controla sus procesos,  
describir las condiciones relacionadas con datos importantes e in-

forma qué decisiones deben tomarse para ser más efectiva,

El viento, las fuerzas de las mareas, el bombeo de agua para usos  
industriales y el agua dulce aportada por los ríos constituyen las  
fuerzas externas mayores que afectan a la bahía aunque cada uno de  
ellos tiene cinco divisiones principales (compartimientos) son únicas en el  
modo de estas fuerzas ejercer su influencia. La precipitación  
pluvial, la evaporación y el drenaje morfoestructural afectan al sistema  
hidrológico en un grado menor. La topografía submarina y la bathi-  
metría de la bahía son los factores más efectivos en el control del  
movimiento de agua en la bahía, siendo características de natura-  
leza única en los compartimientos que componen la misma.

La estructura biológica afecta al hidrociclo de la bahía mediante  
el control de las tasas de intercambio entre sus distintas secciones

© compartimientos. El manclar es importante en los compartimientos  
ennareados por la caleta terna y el sureste de la bahía. Las pre:  
@eras de fanerógamas son importantes en las secciones delimitadas  
por el oeste, sureste y la parte central de la bahía donde Getas  
afectan la profundidad del umbral de la bahía y el material  
acarreado por los ventisqueros.

Los flujos de agua resultantes de la tala de anotonamiento por los  
ventisqueros y retorno al equilibrio son los dos tipos de grandes  
de corrientes que a su vez están más sujetos a inconsistencias ©  
vagueadas. De acuerdo a este estudio, sería beneficioso llevar @  
cabo una investigación para caracterizar las corrientes provenientes  
de anotonamiento de agua por los ventisqueros a través del  
geobanfas Uras junto a una evaluación de flujos resultantes de  
contracorrientes de equilibrio en canales profundos. De esta  
manera se podría predecir el comportamiento físico de la Bahía de  
Guayanitita,

Las opciones 0 alternativas de manejo que afectarían a la bahía podrían muy bien influenciar en su productividad, su uso como puerto y las características del agua a usarse para enfriamiento en las centrales generatrices. Su geomorfología y la velocidad del viento ejercen

un mayor control sobre la temperatura del agua a usarse para enfriamiento. Los umbrales entre los distintos compartimientos de la

bahía y el agua dulce de superficie y niveles freáticos de las cuencas pluviales de los fos Yauco y Guayanilla tienen un mayor control sobre la productividad biológica (desde el punto de vista hidrológico) «

El desarrollo de puertos podría ser afectado por corrientes marinas

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Y comunidades biológicas que controlen las tasas de sedimentación

y estabilicen la topografía submarina. Se articularían las decisiones

y el desarrollo de la bahía podrían ser influenciadas por

la información generada por un modelo hidrológico de manera que se

puedan sostener y aumentar sus usos productivos de una manera

eficiente en cuanto al costo.

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## HYDROLOGIC MODEL, OF GUAYANILLA ERY

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The hydrologic flux among geographic areas is an important process mediating the transfer of energy and materials in the coastal zone. In the coastal bays of southern Puerto Rico this flux is affected primarily by wind drift, tides, runoff (Xunel and Iadjathwodorou, 1970; Golénan, 1979), industrial cooling water pumping, and groundwater flow. Storm Surges and ship traffic affect the water budget less frequently or on a smaller scale

This paper describes a hydrologic model of Guayanilla Bay, Puerto Rico (Fig. 2). The model accounts for the mass balance of water among five major hydrologic subdivisions (or compartments) of the bay, describing daily flux as an average of events that occur on an hourly, daily, monthly, and seasonal basis. It is useful for understanding the relative importance of different flows, their respective controls, and their effect on habitat types and industrial uses within the bay. The model is based on the information summarized below and appended, and should provide a reference for acquiring improved data for more accurate prediction and management applications. This model is one part of a series of models that describe environmental and economic processes in Guayanilla Bay and its surroundings that are formulated at the Marine Ecology Division, Center for Energy and Environment Research, Mayaguez, Puerto Rico.

Guayanilla Bay is located on the south coast of Puerto Rico (Fig. 1) and consists of the five sub-areas (compartments) shown in Fig. 2 and characterized in Table 1. Compartment boundaries are defined by submerged bars, jetties, headlands, seagrass beds, cays and dredge spoil banks. Sills shallower than 0.7 m are indicated in Fig. 2 and are critical for isolating several bay compartments. Openings, including channels and

pipes or other industrial structures, facilitate communication between the compartments. This communication or "interface" between compartments, occurs along the exposure between subareas, and is in part dependent on depth characteristics (see Table 1 and Section 2) +

## 2.0 SOURCES OF WATER MOVEMENT

Five primary factors affect water movement in the bay: tides, wind

at rate, runoff, groundwater, and pumping of industrial cooling water.

Additionally, shipping and storm surges can occasionally affect coastal

water flow. Quantification of these parameters in the following sections

provides the basis for developing the generalized annual hydrologic budget,

as presented in this paper.

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## 2.1 Tapes

The tidal range at Guayanilla Bay varies between 15 to 45 cm

(ge, 1972) with a daily average range of 30 cm.) Based on this range,

£10x for the bay is  $2.65 \times 10^9$  m<sup>3</sup> daily. Tidal flux is a small proportion

Of Central Bay volun: (78). However, tidal flux is a large proportion

Of the water volume of the Western Bay (20+), and is significant in

the other shallowwater compartments. The Western Bay may be the most

"powerful" in terms of productivity and respiration (Chartock, in pre

paration) so that the exchange of materials through tidal forces is

critical for the bay system.

## 2.2 WIND DRIFT

Surface water movement in Guayanilla Bay has a significant influence

from the wind (Goldman, 1979). Surface wind drift has been reported as

2.6 to 5.8 % of wind velocity based on a review of 16 coastal and oceanic

studies (Lange and Wuhnorfuss, 1979). In shallow bays with limited fetch,

Goldman estimates that drift in the upper (2-3 meters) mixed layer is

approximately 1% in South Coastal Caribbean Bays, as supported by studies

during 1977 through 1979 (Goldman, 1979). This is in near agreement with

estimates of 2-34 for fetches of 4 to 10 km summarized by Von AEX (1960),

Estimates of surface water movement induced by wind at Guayanilla Bay

tabulated in Table 2 are based on average 24 hr. wind velocity of 2-9 m/sec

(see Appendix A). The average vector of the wind is easterly, shifting

from the southeast (120°) to the northeast (60°) on a diurnal basis.

Water movement at the Caribbean interface is affected by circadian and seasonal variation in wind direction. This effect is due to the southern exposure of the Central Bay - Caribbean interface and results in switching the wind drift in and out of the bay. This switching has not yet been verified with drift bottle releases (Goldman, 1973). Thus, wind vector fluctuations must be accounted for to estimate flows. As summarized in Appendix 4, the directions of wind predominate: northeast ( $60^\circ$ ) at night, and most of the day during winter months (January through March) and southeast ( $120^\circ$ ) during the day most of the year. The northeast component is dominant approximately 50% of the time, and the southeast component is dominant 60% of the time. These two average wind vectors are included in the wind drift data summarized in Table 2. As indicated in Table 2, the volume of flow from wind drift is most significant for the exposed Central Bay where surface water drifts into the bay during the day and exits at night.

The wind driven flow entering the bay is  $2.92 \times 10^8 \text{ m}^3$  and the flow driven out the bay is  $1.95 \times 10^6 \text{ m}^3$  daily. These quantities compare with the  $2.65 \times 10^8 \text{ m}^3$  per day moving by tidal forces (see Section 3-0), How

ever, the wind driven circulation of the bay is a critical variable for the exchange of water in Guayanilla Bay since this factor fluctuates seasonally and daily, and may control upwelling (see Section 2.6)

National Oceanic and Atmospheric Administration data from Ponce. Confirmed with selected monthly measurements at Guayanilla Bay (Chartock, in preparation) -

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TABLE 2

hydrologic Flux Summary®

(108 n? per day)

Compartment Wind ngustrial Wano®® and

Compartment Vortine ?tidal {Summary) Puncing Groundvater

Western Bay 3.30 og 1.82 - 0.072

central pay 37.6 2.65% 3.89 - 0.150

Intake 1.38 oma a3 te 008

Therma cove 0.52 oor 67216 002

Southeast 1.56 1246 4.08, - -

Exbayent

?total 44.33 aes = - 202

Data tabulated are aggregate flow. Separate flows among compartments:

are presented in Table 5.

b, The tote {lux of the bay system passes through the Central Bay

(volume of Central Bay changes  $1.30 * 10^6$  m) uring s tial cycle).

ce Mind drift estimate of deeper channels based upon movenont of upper

Sn of water at 1% of wind velocity (2.9 a/sec). For shoals or sills

where this depth was less than 3 m average sill depth was used, See Appendix

B for detailed wing drift aca.

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### 2.3 INDUSTRIAL PuMPING.

Hydrologic budgets of the intake, thermal cove, and southeast en  
boynont are affected uignificantly by the Costa del Sur Ofi-fired Power  
Plant, which operates by once=through sea water cooling. iien all six  
units? (boilers) are on-line, the cooling water flow is 37.6 m/sece

Tais plant operates with a power factor of approximately Got so that  
estimated pumping is  $2.16 \times 10^9$  m<sup>3</sup> per day. This pumping alone exchanges  
the volune of the thermal cove five times daily.

### 2.4 RUNOFF AND GROUND

R

The Western Bay receives runoff from the Yauco River Watershed, and the Central Bay receives runoff from the Cuayanilla and Hacaña, RIVET. The Yauco River is impounded at the Luchetti Reservoir, and some flow is diverted out of the watershed. The river valleys in all the watersheds in Fig. 2 are developed for irrigated agriculture, primarily sugar cane. Runoff is highest in June through November, and low from December through May. The average annual streamflow for the Guayanilla River, measured approximately two km upstream from the bay discharge, is  $29.6 \times 10^6 \text{ m}^3$  per year; for the Yauco River is  $13.5 \times 10^6 \text{ m}^3$  per year and for the Hacaña River is  $9.8 \times 10^6 \text{ m}^3$  per year (Crooks, et al. 1968). Summary data for watershed flows are included in Appendix B. Potential water entering bay compartments from groundwater and runoff is  $26.3 \times 10^6 \text{ m}^3$  per year to the Western Bay and  $54.3 \times 10^6 \text{ m}^3$  per year to the Central Bay. Actual freshwater entering is likely to be somewhat less due to domestic freshwater use and evaporation in agriculture:

The geological structure of Guayanilla Bay is heterogeneous with a karst topography in Miocene Limestone that outcrops at the surface (Worelock, et al., 1979). A variety of marine and alluvial sediments occur at the surface. Both the Limestone and unconsolidated sediments



serve as aquifers, and the alluvium has been extensively developed for irrigation supply.

Groundwater influx into the bay has been estimated from average groundwater flow along the south coast of Puerto Rico. A daily average discharge of 2.1 m<sup>3</sup> per linear foot of shoreline has been estimated (Puerto Rico Water Resources Authority, 1972). This estimated value is within a factor of 0.5 of the groundwater flow from the watersheds above Guayanilla Bay estimated by the U.S. Geological Survey (Crooks, 1968)

The freshwater flow is small as an annual average, approximately 8.08 of the tidal flux, but seasonal variation and individual storm events can make this a very important factor, with some measurements of peak runoff of  $1.4 \times 10^6$  m per day for the Guayanilla and Yauco Rivers

Rises in water table are caused by evapotranspiration losses of water entering the lower basin, based on similar losses in U.S.G-3. upper basin measurements.

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(almost one-half the tidal flux during high rainfall periods.) The freshwater flux is also important for maintaining the brackish water conditions of the Western Bay, and this is a critical physical factor in structuring the biotic community and its energy flow patterns.

## 2.5 OTHER SOURCES

Four other sources of variclegic f1ur occur: precipitation, evacor  
Bion, storm surges (storm tides), channel dredging, end ship traffic,  
?These sources of flux, however, are either small of infrequent. Direct  
Precipitasion from the annus average rainfall of 30 cm in the rClatively  
arid coastal environnant (Cintrén, tugo, Pool ana sorriss, 1978) results

in an addition of 7,409 mi of water annusly, divided among bay compartments  
according to surface area. Most of this addition occurs from May through  
October. This flux is abcut three orders of macnitude smaller than  
categories shown in Table 2,

Aznual pan evaporation is 19.9 inchwe at Ponce, with very sinilar  
Values at Guanica and Sabana Grande (Seat, National Occanse end Atmos~  
pheric Administration, personal communication tlovensur, 1379). thin  
Eerults in a flux of  $18 \times 10^9$  md annually cr aperoaimavely 49,000 mi  
Gaily from the entire bay (see Table 1)

Storm surges accompany the tropical depressions, storms, and hurricanes that frequent the Caribbean. The hurricane force winds and storm surge contacts Puerto Rico an average of once every six years (Puerto Rico Water Resources Authority, 1972). These storms have different intensities, but a maximum expected storm surge along the coast would result in about 3m storm tide (Puerto Rico Water Resources Authority, 1972). These tidal surges would pass during a two to four hour period and result in a displacement of  $26 \times 10^9 \text{ m}^3$  about one-half the volume of the bay.

The large container ships and tankers entering Guayanilla Bay displace approximately 100,000 metric tons, or 48,000  $\text{m}^3$ . An average of one tanker enters and leaves the bay daily (Puerto Rico Port Authority Staff, personal communication, 1980). This exchange occurs at the interface between the Central Bay and the coastal water. Thus, Water movement due to tankers traffic is similar in magnitude to evaporation water flux, although very localized.

?Wish Greater rainfall Occurs during tropical storms. Uncontrolled runoff from the three drainage basins can be equivalent to  $1.9 \times 10^7$

(10 om of rainfatt over a 24 hour period in the drainage basin). This

As about 0.4 tines the volune of the bay. Much of this runoff would

be released within a one day period (Gragg Morris, personal communication) «

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Water Flux from Precipitation and Evaporation

(003 m per dey!

Direct

Location Precipitation Evaporation

Western Bay 2. 16.9

central bay 10.60) 23.9

Intake 0.94 an

?Thermal cove 0.58 nat

Southeast Embayment 2.00 4.49

Total 20.7 49.00

10

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Propeller pumping by 2,000 horsepower tugboats displaces water across the bay - coastal water interface. ?The tusboats have a thrust that moves 29,900 m<sup>3</sup> per minute. Approximately four tusboats per day transit the mouth of the bay, each crossing in an avorage of five seconds ?over the Central Bay - coastal interface, This results in an estimated Localized movenent of 10,000 m<sup>3</sup> of water in each direction.

## 2.6 EQUILIBRATING FLOWS

Equilibrating flows are established in the bay that maintain the volume of compartments. Equilibrating flows are the result of gravitational force that results in flow to establish a uniform (level) geopotential surface. For example industrial pumping reduces the volume of the Intake Bay so that water flows into the Intake Bay from the Central Bay to re-establish equilibrium. In this case, surface water of the Intake Bay is moved by wind into the Central Bay, and an equilibrating flow is the cool bottom water from the Central Bay. This movement of bottom water has been substantiated by Grogue and temperature studies (Golénan, 1979). The size of equilibrating flows are mass balance estimates of counter currents. They are based on the assumption that the average daily volumes of the Guayanilla Bay compartments are constant, (see Section 3).

Increased easterly winds force surface water into the Western Bay.

A bottom equilibrating current from the Western Bay is established as a counterflow that exits a narrow channel near Punta Verraco, resulting in an outflow of turbid Western Bay water into the Central Bay. This flow is substantiated by observations of the extension of turbid Western Bay water that moves east into the Central Bay and then south along Punta Verraco.

### 3.0 MODEL AND PROPERTIES

The model of the bay storages and flows is shown in schematic form in Fig. 4, indicating the inputs and discharges of water and the flows among compartments. The external energy sources are listed in Table 4 with a summary description of the magnitude and type of force described in the previous section. The flows are listed in Table 5, including the origin or source compartment when the source of flow is within the Guayanilla Bay system. The transfer coefficient, or proportion of the source compartment that flows each day, is also provided. The largest flow is an upwelling equilibrating counter current flowing into the Intake Bay, largely the result of westward wind drift and industrial pumping. Generally, wind driven currents and equilibrating flows are the largest flows between compartments.

The exchange of water in and out of the system as a whole, however, is dominated by both tidal flow and wind. Wind drives about 108 scfm of water into the bay (J4) than tide, but tides flush much more water from the bay (31) than does wind (93). The magnitude of the equilibrating flow out of the bay is directly related to wind velocity.



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Propeller pumping by 2,000 horsepower tugboats displaces water across the bay \* coastal water interface. The tugboats have a thrust that moves 29,900 m<sup>3</sup> per minute. Approximately four tugboats per day transit the mouth of the bay, each crossing in an average of five seconds over the Central Bay ~ coastal interface. This results in an estimated localized movement of 10,000 m<sup>3</sup> of water in each direction,

## 2.6 EQUILIBRATING FLOWS

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estimates of counter currents. They are based on the assumption that the average daily volumes of the Guayanilla bay compartments are constant (see Section 3).

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A bottom equilibrating current from the Western Bay is established as a counterflow that exits a narrow channel near Punta Verraco, resulting in an outflow of turbid Western Bay water into the Central Bay. This flow is substantiated by observations of the extension of turbid Western Bay water that moves east into the Central Bay and then south along Punta Verraco.

### 3.0 MODEL, AND PROPERTIES

The model of the bay storages and flows is shown in schematic form in Fig. 4, indicating the inputs and discharges of water and the flows among compartments. The external energy sources are listed in Table 6 with a summary description of the magnitude and type of force described

in the previous section. The flows are listed in Table 5, including the origin or source compartment when the source of flow is within the Guayanilla Bay system. The transfer coefficient, or proportion of the source compartment that flows each day, is also provided. The largest flow is an upwelling equilibrating counter current flowing into the Intake

Bay, largely the result of westward wind drift and industrial pumping.

Generally, wind driven currents and equilibrating flows are the largest flows between compartments

The exchange of water in and out of the system as a whole, however, is dominated by both tidal flow and wind. Wind drives about 10% more water into the bay (34) than tide, but tides flush much more water from the bay (J1) than does wind (53). The magnitude of the equilibrating flow out of the bay is directly related to wind velocity.

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TABLE 4

External Energy Sources

Energy

source Name Notes

1 Tide 30 em sea

' elevation change

> win average 2.9 m/sec

: Streamflow and  $87 \times 10^6$  m<sup>3</sup> per

Groundwater year from water~

shea

Precipitation 90 en/year

Evaporation 200 en/year

I, Shipping 60,000 m<sup>3</sup>/day

ty Industria? pumping  $2.16 \times 10^8$  m<sup>3</sup>/eay

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?The intereelationship anong causal forces and storages that determine syston behavior is provided in the system equations in Table 6 and shown in the energy circuit diagan in Fig. 5 (Odum, 1971), Appendix C contains the hydrologic model written in BASIC with annual average values used for external energy sources. ?The simulation results are aleo provided. AP

pendix D lists a model that includes 24 hour wind direction shift and

25 hour tidal day. The equations and computer model can be used to predict responses to changes in either the external energy sources or

the physical and biological characteristics that determine the transfer rates within the Guayanilla Bay system.

Examples of biological properties that control system behavior are the mangrove forests and seagrass beds that control the rate of water movement across channels. Mangroves determine channel width and the seagrass beds stabilize bottoms and reduce wind driven currents (Scoffin, 1970).

Many of the physical and biological structures controlling hydraulic flux have been manipulated by dredging, construction of jetties, and by industrial pumping. For example, changing the intensity of industrial pumping gradually alters the degree of upwelling in both the Intake

and Southeast Bays. The effect of decreased pumping on the Southeast

bay would be to substitute coastal water for the surface waters originating in the Intake and thermal cove areas. The amount of water upwelling in the Southeast embayment would nearly double if the industrial pumping ceased. The implication of selected management options for effective

use of the bay is briefly described in Section 5 and in the Executive Summary.

#### 4.0. MODEL VERIFICATION

Much of the data used to develop estimates of system parameters are based on measurement of relatively stable characteristics such as shoreline dimensions and bathymetry. However, many system parameters are not directly measurable, are stochastic (with a high degree of variability), or must be inferred (e.g. the equilibrating flows). One purpose of the model is to identify parameters that critically affect system behavior but that are poorly understood.

Summarized below are selected descriptions of the reliability of model data and some observations that substantiate fundamental interrelationships.

##### 4.1) STORAGE VOLUME

Compartment sizes are based on measurements from National Oceanic and Atmospheric Administration Navigational charts.<sup>1</sup> These have been

?Wational Oceanic and Rtrospheric Administration, June 3, 1978 11th  
Edition. Charts No. 25681 Bahia de Guayanilla and Bahfa de Tallaboa.

16

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TABLE 6

Syston Equations @

eS

1.) = Western Bay (WE) Volume

BTS KTR RT, + RTO, - T,0, + mGt,2, =H

2. Q) = Central Bay (cP) Yolune

2 Raby = ByTWRy = KyTyOy + Katy + MyPy ~ KelgO> + RTO, ~ Kyl\O,

1 Ey = Fabsa \* Mel \* KT, - 570 TO, ~ BOD

TBaRaPs Bur @y \* BYERS ~ by@ ? 150) + Kt, - K,1,0,

7 Intake Embayment (I) Volume

BS Mets MyTeO) + Ty MI) KE, - KI, + HO, ~ KIO

4.0, = Thermal cove (72) Volume

~RebSRe IST Maly TKS = ATW, - TO, = Tyg

5. Southeast Embayment (SES) Volume

BS MGR NES REM RTOS HRT, KT) = KTR,

RRs \* 1h) 440,

a. O= rate of change of 9

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selectively verified by depth soundings of the bay during Fall, 1979.

Field trip checks on the geographic Soundings have been provided by reviewing aerial photographs and by cruising along the shoreline to

verify land - mangrove inundations. All measurements are probably accurate within one percent. Average depths of Lagoon compartments are based on

map measurements of transects in each compartment, with an estimated error range of 2. Overall compartment depth error is within 5%

Verification of water movements is based on drogue, current meter, and dye studies of current velocity. These studies substantiate wind velocity and surface current measurements in open water areas. However, detailed confirmation of wind-current relationships have not been made. Related studies of wind drift indicate variation of 10 to 300 of actual values (Lange and Muhnerfurs, 1979).

The existence of all equilibrium currents has been substantiated with drogue studies (Goldman, 1976). In addition, the general relationship between wind velocity and the magnitude of equilibrium (return) flows is substantiated by observations; for example, of the extension of turbid Western Bay water into the Central Bay. Wind and current velocity relationships are also substantiated by multi-depth drogue observations (Goldman, 1972).

## 4.2 FE

## SSH WATER INPUT

Surface water inputs have annual variation of + 40% of the mean annual flow. Although large yearly variation in runoff occurs, the long term average yearly runoff values used in this report are probably representative within 10-20%.

Groundwater flow data are subject to considerable uncertainty. However, total freshwater influx can be verified independently by calculating dilution of Caribbean water in the bay. Dilution of Western Bay waters occurs to the range of 28 to 33 parts per thousand salinity. Total freshwater flow data used in this report are consistent with this range of salinity given the rates of tidal and wind driven water flux (see Table 7). Salinity stratification and mixing data have not been evaluated, however, and error range is uncertain,

## 5.0 MANAGEMENT IMPLICATIONS AND CONCLUSIONS

Goanitia Bay is a complex, multicompartiment estuary with distinctive



sub-areas. From a physical standpoint, the hydrology is affected by  
Numerous Separate forces that collectively characterize the bay. From

a biological standpoint the bay compartments also function distinctly,

but are highly interdependent (Chartock, in preparation). The hydrologic  
behavior of the bay can be managed to affect its biological properties

and human uses as an "Industrial Marine Ecosystem? (Tilly, 1979) +

1?

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ante 7

Relationship of Salinity to Flushing in Western Bay?

a

Salinity is accumulated

## Western Bay Fraction of Volume of Flushing

(feed Fresh 1,0 Freshwater ?Tame

es

28 186 622,000 5.635

23 1598 \$26,000 7.3067

20 29 430,000 5.97

2 10 334,000 46

32 237,000 33

3 108 143,000 2.0

Central Bay has a salinity of 34.5 parts/thousand,

and that the volume of the Sestumn Bay of  $3.3 \times 10^9$  Bo ena roves

freshwater input is 72,000 m<sup>3</sup> per day

20

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5.1 MANAGEMENT OF

The relative isolation of the Western Bay from the Central Bay can be modified by dredging the shallow sill that separates these compartments, as one example of a management action. The sill depth and width parameters influence the strata of water that enter or leave Bay compartments and thus the source of water for equilibrating return flows. The channel, for example, that penetrates the sill near Punta Verrace provides for the equilibrium flow of deeper, poorly oxygenated water out of the Western Bay to the Central Bay, and may be one factor maintaining the oxygen concentration in the otherwise Western bay (Chertocks preparation) that sustains a small commercial fishery (Cole; 1976). Thus, effects of alternative dredging plans on the transfer coefficients for wind-driven currents and for equilibrium return flows should be evaluated with the models such as the one described here.

Dredging is a continuing process used by industry in Guayantilla Bay to maintain adequate port conditions. Suspended materials in shallow waters moved by the wind drift are a major source of sediment that is transported in the bay. The sediment budget of the bay can be managed to reduce or divert sediment sources to minimize dredging expenditures.

Beneficial biological populations that stabilize the bottom may also be a mechanism to avoid or minimize costly dredging programs. The model described here only provides an initial framework for sediment management. A detailed sediment budget is needed to implement an effective stabilization program,

## 5.2. WATERSHED MANAGEMENT

Groundwater and surface water flows are important for maintaining the biological and physical characteristics of the Western Bay such as community composition, productivity, turbidity, and total particulate matter.

The freshwater flow is dependent on surface and subsurface development of the watershed (U.S.G.S., 1968). For example, increased groundwater pumping coupled with severe drought, can result in sea water intrusion from the bay into the alluvial sediments in the Guayanilla Valley. Development of storage and groundwater recharge capacity in the three

rivers that enter Gcayanilla Say, needs to be evaluated as a mechentom,

fo both naintain continued groundwater use and sustain brackish water conditions in the Western Bay. Periodic floods affect the shoreline and bathymetric characteristics, especially in the Western Bay. Management Of the long term average flows of freshwater can easily be included tm

?an evaluation based on the model prosented here, but the dramatic changes Produced by periodic floods require additional ?model parancters.

## 5.2 INDUSTRIAL INTAKES AND DISCHARGES

Industrial pumping enhances upwelling at the boundary between the Intake and Central Says, and reduces upwolling at the mouth Of the Southeast

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Bay. Although water charactorsstien are medified in the bay aysten by this industria} aps svarion (Ayes, 117%), tw mydrolecce flow bettwer tha trupreut arta veto sai aleng the Goast and central TEE water Se not atfvctrd. This aster uxsnuoge ie exitieal dn affecting

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Spbside of Cuoyanilla bay Un th> aGjacent Tullabos Bay) woud elinioate  
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greatly increas the amount of coastal water ontering the bay wien ton  
Pogential for subsctial change in the ?Eay'y ;hysteal and biologiees  
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rent opt.

One major use of the bay as a resource is to maintain or improve the efficiency of the Costa del Sur power plants and each one meets its own water requirements at the heat sink that circulates heat from the geothermal power plant. Heated water from the geothermal scale southward and naturally returns to the sea, thereby making the system self-sustaining. The combination of surface wind and power, however, alone were not functioning (i.e., there were no windmills), and the power plants were pulling in water that has free flow from the southeast. Thus, even small wind velocities provide

an average of 2 m/s from subsurface waters at an average depth of 100 m, significantly improve power plant efficiency. The estimated cost of installation is in the range of \$5,000 to \$10,000 daily, depending on

fuel costs.)

Because water exchanges are affected by wind and bathymetry, managers of the bay should be careful in adjusting bathymetry or bathymetry? erat mnea ERE Ne Wied regime, especially on the eastern narsine of the bay where wind effects are most closely coupled with man's uses and where wind fetch is short and potentially affected by shoreline configuration.

## 5.5 RESEARCH APPLICATIONS

A significant amount of research has been conducted in Guayantilla Bay to characterize the physical and biological structure of the bay and the consequences of industrialization (Gonzalez, 1979). This model summarizes more and can be used to review this research against the needs of and to explain the function of the bay, and to inform planning and management.



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eters used in this node) are subject

of uncertainty, some are more serious than others to widespread

bay's function. They may affect the magnitude of water

across shallow extent of equilibrium flows. Other important

flows are the surface water, groundwater and the

Some amount of variation in tide level. However, it is likely that

2. Research effort on these latter categories would provide

benefit to decision makers. In contrast, a research project

Selective of understanding wind and water flows would be most

Setivity and result due to variation of the critical variables

of Guayanilla Bay.

Although all of the

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APPENDIX A

Wind Data Supplement

Wind Speed and

Wind speed data used in this report are based on 8 years of observations taken at Santa Zsabel Airsort during the period 1946-1953. These indicated an annual mean wind speed of 3.5 m/sec, a less comprehensive set of data for Tallaboa bay, (adjacent to Gueyanilla Bay) show that winds recorded over the one year period beginning June 1, 1975, and ending May 31, 1976, averaged 208 w/ces, when

erected from the 250 ft measurement level to @ level 19 m above the ground.' Wind vector data are summarized in Table Re?

Summarizes the basis for wind drift flows between compartments based on direction of exposure, depth, and the vector of wind velocity,

?Fucrto Bice Water Resources Authority, 1976. South Coast Pousr Plane  
Complex. p. 42.

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TABLE At

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Santa Isabel, Purrto Rico

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october, 24 ?

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Based on the period 1946-1953.

Sources: Puerto Rico Water Resources Authority 1976. "South coast

Power Plant Complex? Table 4. 2-1.

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Footnotes to TABLE Ac2

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Channel between Cayo Mata and Commonwealth O12 Refining Corporation (CORCO) Jetty.

Channel between Punta Gotay and Cayo Mata.

Channel between CORCO, Jetty ané Punta Papilio.

Communication from east side of Intake embayment is through forced punping to thermal cove.

Deep portion of channel from Punta Gotay to reef midway to Punta Verraco.

Shallow portion of channel between reef (see @ above) and Punta Verraco.

Sectional area for wind drift is the width times sill depth or 3m, whichever is shallower.

Wind drift calculation is based on velocity of wind driven water through channel. Movement of water is estimated at

18 of average wind speed (2.9 m/sec), and direction (easterly) «

Wind drift occurs in the upper three meters unless a sill is present, in which case the average sill height is used.

Wind @izection has an annual average component of 120° for 60% of the year, and 60° for 40% of the year based on diurnal and seasonal variations in Appendix B (and summarized by the National Weather Service, 1978). To account for component of wind drift through embayment exposure, cosine of velocity component and

time duration corrections have been made for wind vector corrected

@aily flow. For most exposures, this results in a 0.866 correction

factor. For the southern exposure of the Central Bay, the correction factor is 0.2 for flows to the south and 0.3 for flows to the north (e.g. cosine of angle of wind incidence normal to

the exposure of the bay multiplied by duration of the wind).

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## APPENDIX 8

### Surface and Groundwater Data

TABLE 8-1. Annual rainfall on the three principal river basins,

1961 and long-term average.

sn

Drainage Weighted Rainfall

Area cr

River Basin xn? 1961 tong-term

Rio aves

Upper basin, (excluding Lago 39.1 132 163

Lucehetti diversion)

Lower basin %

Entire basin, (excluding Lago 101 127

Lucchetti)

fo Guayanst1a

Upper basin, above stream station tes 196

Lower basin 8 9

Entire basin 145 160

Rio macané

Upper basin 19.5 198 15

Lower basin 33.7 36 6

Entire basin 332 150 150

?Three basins, (

?Euccherss)

Upper basins 50.2 168 180

Lower basins, the Guayanilla~

99

145

Source: Modified from Crooks et al. 1968

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TABLE B+?

Annual amount of water received by the Cuayanilla-Yauco River basins,

4961-63 and long-term average.

amount, 10° m?

Source of water a

iverase

Rainfall on upper basing

Rio Yauco (does not include 82 a) os

Grainage area above Lvccheets

ane)

Rio Guayanitla 89 m4 9

Rio Macand 31 B38

?Total rainfall on upper basins 178 151236 191

Streamflow entering lower basins

Rio Yauco (above first diversions) 20 n 16 4

Rio Guayaniilla (gaging station) 27 " 2 2

Ro macan ace £ 2 10

?Total streamflow entering 58 OST 33

ower basins

Rainfall on tower basins

Rio Yauco 2 2 oe 30

rio Guayaniite 23 2 2 2

Rio Macand 42 BON 15

ee so @

Water reaching lower basins 12 971s 135

Source.

Modified from crooks et al. 1968

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Listing of Computer Program (in BASIC) with Energy Sources Constant

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Bee

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A HYDROLOGIC MODEL OF GUAYANILLA BAYPRe

By

MICHAEL A. CHARTOCK,

CENTER FOR ENERGY & ENVIRONMENT RESEARCH

UNIVERSITY OF PUERTO RICO & U.S. D.O.E.

January 1960

TABLE OF COMPARTMENT VOLUMES

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AY WESTERN BAY CENTRAL BAY INTAKE Bay THERNAL COUE SOUTHEAST BAY

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Limiting of Computer Program with Hourly Tidal and Wind Variation

and Table of Component Values

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432 GRAYT, GUERSTON Br PRINTS OR PLOTS HOURLY VOLUME?

UPAR" 35)

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50 PRINT = MICHAEL Ae CHARTOCK?

360. PRINT

470 PRINT ?CENTER FOR ENERGY & ENVIRONMENT RESEARCH\*

398 ERINT UNIVERSITY oF evento glco a uss. bios."

153 PRINT unm Sakiinn T8062

200 REM REFER TO DAILY VERSION (APPENDIX C) FOR ADDITION NOTATION

349 REM STATEMENTS 490 Of 1600 CONTROL PRINT OR PLOT OPTION

320 REM TO'PLOT CHANCE T600°T0 ?GOTO Good! AND DELETE STATEMENT 490

332 BEN SEGTION 900 CONTROLS HOURLY CHANGES OF uIND AND. TIDt

BRS RED AGHAGLY EGR SPNUEREEALPUEYASHOLES BEAMINPLSNPOTTDE reous on syste  
ne

Eugvrer.

290 REM SEE APPENDED TABLE FOR BEHAVIOR OF PRESENT CONFIGURATION. NOTE  
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QA OF VOLUME OF CENTRAL BAY FROM NORTHEAST WIND IN HORINING FOLLOVED BY  
FURTH

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420 02 = 3746

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46 8S = 9:82

439 G9r0 690

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710 LET Ka = (2092) 7 1

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720 LET Ky = C0821 7 5.3007 T

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731 REM ?The FROM CR TO UB

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RAINFALL INTO WB

$K_e = (0.15) T$

FRESHWATER TO CR

$K_D = (0.24 / 37.6) T$

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Cate tewas

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875 LET US = 1.3767 7 37.8) 7 T

876 REN COUNTERCURENT FROM CR? TO CARTE

989 REN GRAPHICS SECTION @

900 REN CALCULATES STORAGE VALUES

903 LET He = HR +a

wo Tan = $\phi$

15 TF WR = 25,

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920 IF ww > 12

THEN 12 =



925 IF Wu < 13 THEN 12 =

928 TE WW = 24 THEN Hu

930 LET'Iq ad

950 LET 15 = 4

9e0 Ler ie = i

979, Ler 17 = 3

680 ?Rew Tine surrch

4010 IF 11 = > 0 THEN Ki = Ofk7 = OtkK = orkR = o:kY = 0

4020 IF 11 < 0 THEN K2 = O:KB = OK = OK = OIKU =o

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32.805) ~ CRU #11 #02) + (KU # IL x OS) - (LO ¥ 02) - (LE HOD) + CKY ETE

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3) + (KZ x 02) = (KL £7 4 03) + 03

17803) ~ CH #15 G4) + (KO HIS) + (KP Tt 4 a5) ~ CKO

Ey oGSR A 32 £ ?Cua tan + a4

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Th #02) ~ (KY BTL x05) ~ (KU E12 #05) + C42 a a2d ECL

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TABLE OF HOURLY COMPARTMENT VOLUKES\*

CNTLLIONS OF CURIC RE

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TAPLE OF MOURLY COMPARTHENT VOLUMES

CHTLLIONS OF CUBIC RETERS

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£ SS700GSE TTFERRTF «TENGE 1168088962  
2 32+2690165 L.443943ge 552336374 1.49949199  
3 Saiissvobs 148311087 © SEBSESE7} 1198818192  
? 37:0354331 1147740203 Isragesgea 1177333081  
3 Serp0ei300 ?T4G74NSSS ELLE 1200308883  
\$:3Q043H3S 5g.7475761 ragszaix7 ? Lseoasgo12 td vasaaeo7a  
\$ HPSS SELAH | NASEZEISG SRERQE? 4d s04n2zs  
8 \$:19240952 34.4651306 1149728335 se7oeso7a \_1.87034393  
9 Si1so747sh ? 3E13032e8° agsaktas? ? Edosebo? 1 1abSStES  
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3 5102203838 -S6logorses © FT2EETSE? TEGSSELASP 1190898  
42 B:99S9073 ? \$6-295342, 1543207134 -Seesea7i s.orsanezs  
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